

DEVELOPMENT OF AN IMPLEMENTATION PLAN FOR ATMOSPHERIC CARBON MONITORING IN CALIFORNIA



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Development of an Implementation Plan for Atmospheric Carbon Monitoring in California

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Abstract

This report describes the design of atmospheric CO₂ concentration measurements that, in combination with other measurements and models, would be used to quantify regionally distributed CO₂ exchanges from California's terrestrial ecosystems and CO₂ emissions from fossil fuel combustion. Using models of net ecosystem CO₂ exchange (NEE), fossil fuel CO₂ emissions, and regional meteorology, we predict CO₂ concentration "signals" in the atmosphere. The predictions of NEE exhibit spatial and temporal variations that are controlled by land cover and climate. Fossil fuel CO₂ emissions from metropolitan areas are the strongest localized sources of CO₂ while weaker but spatially extensive fossil emissions are present throughout the Central Valley. We subdivide the CO₂ sources into four components: NEE inside and outside CA, and fossil fuel CO₂ inside and outside CA. Maps of predicted atmospheric CO₂ concentration signals from these four sources largely mirror the instantaneous emissions near strong sources but plumes of CO₂ enriched or depleted air are predicted to advect far from their sources. We then identify a baseline set of observing stations from existing and possible future sites that could be used to characterize in-state and out-of-state ecosystem and fossil fuel contributions to atmospheric CO₂ concentrations. For each of the stations we calculate mean midday concentration signals with standard deviation for each month and source. We also calculate the covariance of the signal due to NEE inside CA with each of the other signals to quantify how much of the signal from NEE inside CA might be readily separable from the other signals. On the basis of these predictions, we identify new observing stations and a measurement protocol that, in combination with existing stations, would provide data to estimate NEE within CA. Although beyond the scope of this project, future work should estimate the uncertainties in estimating California's NEE that would be obtained using atmospheric concentration data from the stations identified herein.

Keywords: greenhouse gas emissions, atmospheric monitoring, inverse methods, land-surface CO₂ modeling

Preface

The Public Interest Energy Research (PIER) Program supports public interest energy research and development that will help improve the quality of life in California by bringing environmentally safe, affordable, and reliable energy services and products to the marketplace. The PIER Program, managed by the Energy Commission, annually awards up to \$62 million to conduct the most promising public interest energy research by partnering with Research, Development, and Demonstration (RD&D) organizations, including individuals, businesses, utilities, and public or private research institutions. PIER funding efforts are focused on the following six RD&D program areas:

- ▶ Buildings Energy Efficiency End Use
- ▶ Industrial/Agricultural/Water End-Use Energy Efficiency
- ▶ Renewable Energy
- ▶ Environmentally-Preferred Advanced Generation
- ▶ Energy-Related Environmental Research
- ▶ Energy Systems Integration.

What follows is the draft report for the PIER-EA Exploratory Grant contract, Contract Number #500-02-004 – MRA #015-009, conducted by the Lawrence Berkeley National Laboratory. The report is entitled “Development of an Implementation Plan for Atmospheric Carbon Monitoring in California” This project contributes to the Energy-Related Environmental Research program. For more information on the PIER Program, please visit the Energy Commission’s web site at <http://www.energy.ca.gov/research/index.html> or contact the Energy Commission’s Publications Unit at 916-654-5200.

Executive Summary

The Issue

Carbon dioxide (CO₂) is a greenhouse gas (GHG) that is emitted by fossil fuel combustion for electricity production, transportation, industry, and other human uses. Globally, roughly half of the CO₂ emitted by fossil fuel combustion remains in the atmosphere, leading to sustained increases in atmospheric CO₂ concentrations. In-state electricity generation accounts for about 16% of California's anthropogenic CO₂ emissions.¹ Because these emissions contribute to human-induced climate change, the California Energy Commission (Energy Commission) has identified the study of CO₂ and other GHGs as a priority area for research, with a particular focus on identifying CO₂ sources. Future cost and production strategies for electricity generation may be affected by economic and policy responses to global warming. Therefore, it will be important to determine the contribution of fossil fuel combustion and other sources of CO₂ exchange to atmospheric CO₂ levels.

Currently, the Energy Commission estimates CO₂ emissions using accounting data for fossil fuel emissions and a combination of inventory surveys and simple models for net ecosystem exchange² (i.e., the net balance of ecosystem carbon fluxes into and out of the atmosphere, NEE) of CO₂. The net ecosystem fluxes occur continuously and uncertainties associated with these fluxes are poorly quantified.

At the national level, the carbon cycle community was posed with the analogous problem of providing defensible estimates of net CO₂ (and other GHG) exchange. In response, the research community suggested a North American Carbon Program (NACP) to provide a robust approach for reducing the uncertainties in regional CO₂ exchange estimates. The NACP is envisioned to combine inventory methods of biosphere carbon stocks and fossil fuel emissions, direct CO₂ flux measurements³, and measurements of atmospheric CO₂ concentrations, all coupled within an inverse analysis⁴ employing ecosystem-atmosphere models of NEE and atmospheric transport.

¹ California Energy Commission. November 2001. *Inventory of California Greenhouse Gas Emissions and Sinks: 1990–1999*. P600-02-001F. Figure ES-8.

² NEE is the net flow of CO₂ from terrestrial ecosystems to the atmosphere. For both NEE and fossil fuel CO₂ emissions, fluxes are defined so that a positive flux causes an increase in atmospheric CO₂ concentrations.

³ Direct flux measurements quantify CO₂ exchange over a limited area using micrometeorological methods. For example, the eddy covariance method employs rapid measurements of vertical wind velocity and CO₂ concentration to calculate the net flux of CO₂ as average of CO₂ concentration weighted by vertical velocity. Equivalently, the flux can be viewed as the net difference in the amount of CO₂ carried upward minus downward with the atmospheric eddies.

⁴ Mathematical inverse methods estimate the spatial and temporal distribution of surface fluxes that are most consistent with measurements of CO₂ concentration by working backward from concentration to flux under the assumption that CO₂ concentrations are the result of surface fluxes being transported in the atmosphere by

Measurement of CO₂ concentration gradients on regional scales is challenging because the concentration differences generated by NEE are in the range of less than 1 ppm to approximately 20 ppm. Sufficiently accurate measurements can be accomplished for signals greater than about 1 ppm from the instrument perspective, as the state of the art in measurement accuracy is approximately 0.2 ppm of CO₂. A careful examination of the expected concentration signals from NEE and fossil fuel CO₂ emissions is an important first step in assuring that atmospheric methods can be applied successfully to California.

Anticipated Benefits for California

This work benefits California electricity rate-payers by providing the foundation for a network of atmospheric monitoring sites and trace-gas sampling protocols that will enable researchers to more precisely identify the CO₂ contributions from ecosystem processes and fossil fuel combustion to global warming. The resulting tool will enable the quantification and verification of net CO₂ emissions, strategies for within-state carbon sequestration, and emissions trading. This project addresses state and national needs for quantification of terrestrial sources and sinks of carbon cycle gases.

Project Objectives

With the goal of designing a system of atmospheric concentration measurements as part of a larger strategy for monitoring CO₂ exchange in California, we identified four objectives:

- I. Predict temporal and spatial distributions of NEE and meteorology in the Western US for representative time periods covering the seasonal cycle of NEE.
- II. Predict temporal and spatial distributions of fossil fuel CO₂ emissions in the Western US.
- III. Predict temporal and spatial distributions of CO₂ concentration signals due to NEE and fossil fuel CO₂ emissions inside and outside California.
- IV. Use the atmospheric CO₂ concentration signals predicted for existing and potential future monitoring stations to judge the suitability of these stations for quantifying NEE within CA.

meteorology. See for example, Kasibhatla, P. and American Geophysical Union (2000). Inverse methods in global biogeochemical cycles. Washington, DC, American Geophysical Union.

Key Results

- The predicted atmospheric CO₂ concentration signals from NEE at many locations within CA are both measurable with existing methods and often comparable or significantly larger than signals from fossil fuel emissions.
- The predicted CO₂ concentrations at marine buoys and coastal stations are significantly affected by terrestrial signals and cannot be used to represent global background CO₂ concentrations without thoughtful correction.
- A combination of existing terrestrial sites including flux measurement sites and tall communication towers would likely be effective as observation stations for separating CO₂ concentration signals from NEE within CA from signals due to fossil fuel emissions and NEE outside CA.

Project recommendations

CO₂ observing stations, sampling frequencies, measurement protocols:

- New terrestrial sites sensitive to California NEE: New terrestrial observation stations should be deployed, likely on existing communications tall towers in California where signals are predominantly influenced by NEE or fossil fuel CO₂ emitted within CA.
- Sampling frequency: Air sampling should be conducted continuously in an automated manner to accurately capture the concentration during the well-mixed afternoon period and the diurnal variation. Data should be collected and averaged into approximately 30 minute intervals for further analysis. However, if resources are limited, periodic sampling with National Oceanic and Atmospheric Administration's (NOAA) flask sampling systems⁵ could capture day versus night differences, providing valuable initial data to test the model predictions from this and other studies.
- Measurement protocol: Air sampling and analysis, instrument design, calibration, and analysis should follow established methods to ensure accuracy and minimize systematic errors. Continuous measurements should be checked against measurements with NOAA flask systems. In addition to the CO₂ data, standard ancillary meteorological data (e.g., temperature, relative humidity, wind velocity) should also be gathered and recorded.
- Additional stations: Observations are planned for a variety of platforms to monitor background air entering California for carbon cycle and air quality purposes. This will likely include the three planned marine stations planned by NOAA, periodic aircraft profiling planned by NOAA, and sampling from one or more mountain-top stations. It is also expected that existing flux measurement sites will be augmented with accurate and

⁵ NOAA conducts routine sampling and laboratory analysis of air samples using standardized methods. <http://www.cmdl.noaa.gov/ccgg/flask.html>

precise CO₂ concentration measurements. These data sets should be included in future analyses.

Uncertainty Analysis and Station Optimization

- Future work should expand on the present study to predict the uncertainty that will likely be obtained in estimating the total NEE from inside California from the measured concentration data.
- Each of the major sources of uncertainty (e.g., spatial coverage of stations, background subtraction, and transport error) should be quantified in terms of uncertainty in estimating NEE.
- Future simulations should include several complete annual cycles to sample different weather patterns and to capture inter-annual variations in NEE.

Measurement of additional trace gas species:

- Where possible, carbon monoxide (CO) measurements should be included with CO₂ measurements to quantify fossil fuel CO₂ emissions. The use of other sources of CO (and other pollutant) data from the California Air Resources Board (CARB) might play a valuable part in this effort, particularly near urban areas where pollutant concentrations are high enough to be clearly detected with available instruments.
- Because fossil fuel CO₂ contains no radiocarbon (¹⁴C), measurements of ¹⁴CO₂ in atmospheric samples can be used to infer the fraction of CO₂ due to fossil fuel combustion. Because of the high cost for precision analysis, this technique can probably only be applied with a low sampling frequency.
- The use of ²²²Rn as a tracer of atmospheric mixing should be explored for use in combination with CO₂ measurements to quantify net regional CO₂. This technique may be well suited to the problem of CA carbon budgets, because much of the air entering CA is of marine origin and will be depleted in ²²²Rn.

Data Synthesis:

- A crucial aspect of monitoring California's CO₂ emissions using the concentration measurements in an inverse or "top-down" approach will be to sensibly incorporate "bottom-up" information (e.g., ecosystem production from forest inventories and crop yields, fossil CO₂ emissions from improved inventory methods, eddy covariance CO₂ flux data from towers and aircraft flights, and ecosystem models of CO₂ exchange). This approach is an emerging area and developments in the North American Carbon Program (NACP) will be valuable to addressing these issues in CA.

Estimating Non-CO₂ GHG emissions:

- The methods that we have applied here could be used to develop a measurement strategy for non-CO₂ GHGs. Because the relative uncertainties for fluxes of non-CO₂ GHGs are considerably larger than for CO₂, atmospheric budget and inverse approaches may be particularly effective in significantly reducing those uncertainties if atmospheric observations can be performed.

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1. Introduction

1.1 Background and Overview

The Intergovernmental Panel on Climate Change (IPCC) concluded that anthropogenic increases in atmospheric GHG concentrations are affecting global climate (IPCC 2001). In California and the Western U.S., recent work indicates that the future quality of environmental services may be compromised by changing climate (Wilkinson et al. 2002; Hayhoe et al. 2004). In response, the Energy Commission has identified quantification of California's net GHG emissions as a research priority. As such, the dominant sources of CO₂ emissions from fossil fuel combustion for electricity generation and transportation will be considered in the context of California's total GHG balance. This balance includes contributions from fossil fuel emissions and terrestrial NEE as affected by climate and land-use.

NEE is difficult to quantify on regional scales because (a) the drivers (e.g., climate, land use, and management) are hard to quantify; (b) estimates of gross respiratory and photosynthetic CO₂ exchanges are uncertain; and (c) measurements of these exchanges are difficult and expensive to make. Currently, the Energy Commission estimates CO₂ emissions using accounting data for fossil fuel emissions and a combination of inventory surveys and empirical models for NEE (Franco 2002). Inventory and simple budget methods are in place to estimate some stocks of ecosystem carbon pools (e.g., aboveground plant biomass, litter, and soil organic carbon) in some locations. However, it is difficult to quantify statewide carbon sources or sinks from changes in stocks because it is expensive to study large areas in sufficient detail and because large uncertainties result from calculating NEE as the difference of large and uncertain stocks. For example, recent work found large inconsistencies between carbon fluxes estimated using US Department of Forestry data in Franco (2002) and fluxes estimated using data provided by the California Department of Forestry (Winrock 2004).

At the national level, the carbon cycle community was posed with the same problem of providing defensible estimates of net CO₂ (and other GHG) exchange. In response, the research community suggested the North American Carbon Program (NACP) to provide a robust approach to estimate exchange of CO₂ and other carbon cycle GHGs on regional to continental scales (Wofsy et al. 2002) (<http://www.esig.ucar.edu/nacp/>). Similar efforts are underway in Europe (<http://www.bgc-jena.mpg.de/public/carboeur/>). The NACP is envisioned to combine inventory methods of biosphere carbon stocks and fossil fuel emissions, direct CO₂ flux measurements, and atmospheric monitoring of CO₂ concentrations, all coupled with inverse analysis employing ecosystem-atmosphere models of NEE and atmospheric transport.

Atmospheric measurements facilitate independent and complimentary approaches to estimating NEE. Micrometeorological methods (e.g., eddy covariance flux) are now employed in networks of ground-based towers to directly measure long-term NEE on plot scales (~ 1-10 km) (Baldocchi et al. 2001) and aircraft-based platforms to directly measure short-term NEE on regional scales (10-100 km) (Desjardins et al. 1997). At present, techniques for estimation of seasonal to inter-annual NEE on larger regional scales (~ 100-1000 km) are being developed. Some of these include mixing models using a tracer with a known source emission rate (Biraud et al. 2000; Biraud et al. 2002) and atmospheric boundary layer budgets based on assumptions of

average dynamics in the atmospheric boundary layer (Lloyd et al. 2001; Bakwin et al. 2004; Helliker et al. 2004). As part of NACP, ambitious attempts are now underway to refine inverse techniques originally developed to estimate NEE on global to continental scales (Ciais et al. 1995; Denning et al. 1996; Gurney et al. 2002).

A key feature of the inverse approach is to use NEE predicted by land surface models as an *a priori* estimate of NEE which is then refined so as to minimize the difference between predicted and observed atmospheric CO₂ concentrations to produce an *a posteriori* estimate of NEE. This approach requires a model of atmospheric transport to relate surface fluxes of CO₂ to atmospheric concentrations of CO₂. The value of this approach is that the uncertainties in the *a priori* NEE, the transport model, and measured CO₂ fluxes and concentrations can be propagated through the analysis to estimate uncertainty in the *a posteriori* NEE estimate. By comparing the *a priori* and *a posteriori* uncertainties in NEE one can judge whether the atmospheric inversion approach has been effective in improving (reducing uncertainty of) the NEE estimate.

Several future sources of atmospheric data will aid the development of regional estimates. First, the NOAA's Climate Monitoring and Diagnostics Laboratory (NOAA-CMDL) has plans to expand their existing operations (<http://www.cmdl.noaa.gov/publications/annrpt27/>) to include additional ground-based sites and vertical profiles of CO₂ and other trace gases using aircraft. Second, in 2003, NASA approved the Orbiting Carbon Observatory (OCO). OCO, planned for a 2007 launch, is designed to provide column-integrated measurements of atmospheric CO₂ with ~ 1 ppm accuracy on 10 km swaths separated by roughly 100 km on a 16 day repeat cycle (Crisp et al. 2004). The global coverage and temporal resolution afforded by OCO will be unique, although the coverage will be sparse at regional scales and the column-integrated signals are significantly smaller than those obtained from surface stations (Olsen et al. 2004). Hence, despite these advances, no current plans include atmospheric CO₂ concentration measurements specifically designed to enable accurate estimation of California's net CO₂ exchange.

Measurement of CO₂ concentration gradients on regional scales is challenging because the spatial gradients are small. In a simplification that aids discussion, the difference in atmospheric concentration ΔCO_2 (ppm) between two towers can be expressed in terms of the average surface rate of net exchange $\langle\text{NEE}\rangle$ ($\mu\text{mol m}^{-2} \text{s}^{-1}$) as

$$\Delta\text{CO}_2 \text{ (ppm)} = \langle\text{NEE}\rangle L / (\rho u h), \quad (1)$$

where L (km) is the tower separation (km), ρ is the molar density of air ($\sim 40 \text{ mol m}^{-3}$), u (m s^{-1}) is the mean wind speed, and h is the height of the mixing layer (km). For example, $\Delta\text{CO}_2 / \langle\text{NEE}\rangle \sim 1 \text{ ppm} / (\mu\text{mol m}^{-2} \text{s}^{-1})$ for a tower separation, $L = 200 \text{ km}$, mean wind, $u = 5 \text{ m s}^{-1}$, and mixed layer height, $h = 1 \text{ km}$. We note that for a tower separation, $L = 200 \text{ km}$, an air parcel moving over the landscape at 5 m s^{-1} has a transit time between towers of 12 hr or $\frac{1}{2}$ day. Since the average uptake over a daytime period is likely to be only -1 to $-10 \mu\text{mol m}^{-2} \text{s}^{-1}$, the differences in CO₂ concentration are 1 to 10 ppm. Measuring the 1 ppm to 10% precision requires relative instrument accuracy of 0.1 ppm. This accuracy can be accomplished from the instrument perspective, as the state of the art in measurement accuracy is approximately 0.1-0.2 ppm (Bakwin et al. 1998). High accuracy is necessary because an average annual flux of just $0.1 \mu\text{mol m}^{-2} \text{s}^{-1}$ is equivalent to a change of $0.38 \text{ tC ha}^{-1} \text{ yr}^{-1}$. Hence, a careful examination of the

expected signal levels is an important first step in determining whether inverse methods could be successfully applied to California.

1.2 Project Goal and Objectives

The goal of this project is to design a set of atmospheric concentration measurements as part of a larger strategy for monitoring CO₂ exchange in California. We identified four objectives to meet this goal:

- Predict temporal and spatial distributions of NEE and meteorology in the Western US for representative time periods covering the seasonal cycle of NEE.
- Predict temporal and spatial distributions of fossil fuel CO₂ emissions in the Western US.
- Predict temporal and spatial distributions of CO₂ concentration signals due to NEE and fossil fuel CO₂ emissions inside and outside California.
- Use predicted atmospheric CO₂ concentration signals to identify existing and potential future monitoring stations for suitability in characterizing and separating NEE within CA from other sources.

1.3 Report Organization

Section 2 of this report describes our approach. Section 3 describes our results, including spatiotemporal NEE distributions (Section 3.1); spatiotemporal fossil fuel CO₂ emission distributions (Section 3.2); spatiotemporal distributions of atmospheric CO₂ concentrations (Section 3.3); and identification and evaluation of observing stations (Section 3.4). Section 4 presents conclusions and recommendations on how the Energy Commission can use the project results to choose stations and measurement protocols to provide the most effective augmentation of other existing and planned atmospheric monitoring efforts.

2. Project Approach

This section describes the methods used to address each of the objectives described above.

2.1 Prediction of NEE and Meteorology

We predict NEE and regional meteorology using a modeling framework (MM5-LSM1) that includes a coupled version of MM5 (Grell et al. 1995) and LSM1 (Bonan 1996). The coupled model was chosen because it provides consistent predictions of NEE and surface energy fluxes (latent and sensible heat) that drive atmospheric mixing. In this section, we describe the coupled model and how it was employed for this work.

MM5 is a mesoscale meteorological model that is used for forecasting regional scale weather and for studying atmospheric dynamics and pollutant dispersion (for a partial list of papers see: <http://www.mmm.ucar.edu/mm5/Publications/mm5-papers.html>). The model can be run with

nested grids so that large-scale atmospheric features can be captured as well as impacts of fine scale topography and surface fluxes. LSM1 is a land-surface model that simulates CO₂, H₂O, and energy fluxes between ecosystems and the atmosphere. Twenty-eight land cover types are simulated in the model, each with a specified seasonality for leaf area and parameters controlling soil CO₂ flux to the atmosphere. Photosynthesis and ecosystem respiration are calculated at each time step using the land cover information and calculated meteorology in each pixel. LSM1 has been tested in a range of ecosystems at the site level (e.g., (Bonan 1995; Bonan et al. 1997; Riley et al. 2003)). The integration of LSM1 with MM5 was accomplished via the established interface for the OSULSM, with changes in the interface to account for partitioning shortwave radiation between diffuse and direct components (Chen et al. 2001). Previous work has shown that the coupled model's predictions of energy fluxes, near-surface air temperatures, and soil moisture and temperature compared favorably to data collected during a previous field campaign (Betts et al. 1998; Cooley et al. 2004).

We performed four simulations covering the first 15 days of March, June, September, and December of 2003. Although a full year was not simulated due to computational constraints, these periods represent a realistic range of meteorological conditions and ecosystem CO₂ exchanges over the year. The model was run with two nested grids with horizontal resolutions of 60 over the entire US, and 20 km over the modeling shown in Figure 1. The vertical resolution in both grids was 18 logarithmically spaced vertical layers between the surface and 20 km altitude. The following physics packages were used in the simulations: Grell convective scheme, simple ice microphysics, MRF ABL scheme, and the CCM2 radiation package (<http://www.mmm.ucar.edu/mm5>). We used the standard initialization procedure for MM5v3.5, which applies first-guess and boundary condition fields interpolated from the National Center for Environmental Prediction (NCEP, <http://www.ncep.noaa.gov/>) reanalysis data (Kalnay et al. 1996; Kistler et al. 2001) to the outer computational grid.

To improve the accuracy of the seasonal cycles of carbon exchange across the different ecosystem types, we compared the predicted NEE with results of eddy covariance flux measurements made in the Western US and available through the FLUXNET data archive (<http://daacl.esd.ornl.gov/FLUXNET/>) and archival literature (Goldstein et al. 2000; Baldocchi et al. 2001; Falge et al. 2002; Law et al. 2002). Because there are relatively few sites for flux measurements compared to the large number of ecosystem types present in the Western U.S., we aggregated the land cover types present in LSM1 into a subset of seven cover types shown in Figure 1. We then calculated the mean NEE for each cover type using the results of our model runs, with the exception of the desert and urban cover types which produced zero NEE in the model. In the cases where the calculated mean NEE is significantly different from the annual mean flux reported for that cover type, we scaled the daytime and nighttime exchange separately by constant factors for each of the four seasons, so that the daytime and nighttime NEE matched the diurnal cycle from the measurements and the annual mean NEE matched the FLUXNET estimates. Our approach is similar to previous work where the prediction of photosynthesis is assumed to be correct but soil respiration of CO₂ is scaled by a constant factor in each pixel to produce zero net annual CO₂ exchange (Denning et al. 1996). Our method differs from that approach in that we used field measurements to determine the mean annual carbon balance in a given cover type and did not require that individual pixels were in carbon balance. We believe this is appropriate because many California ecosystems are managed to accomplish a net

removal of CO₂ from the atmosphere (e.g., timber forestry and crop agriculture). In our approach we calculated the annual mean NEE and also the mean NEE for daytime and nighttime in each season from the measurements. If the model predictions of mean annual NEE averaged over a given cover type was significantly different from the mean calculated from field measurements then we scaled NEE in all pixels of that cover type by a constant factor so that both the annual mean and the mean daytime and nighttime NEE matched the field measurements.

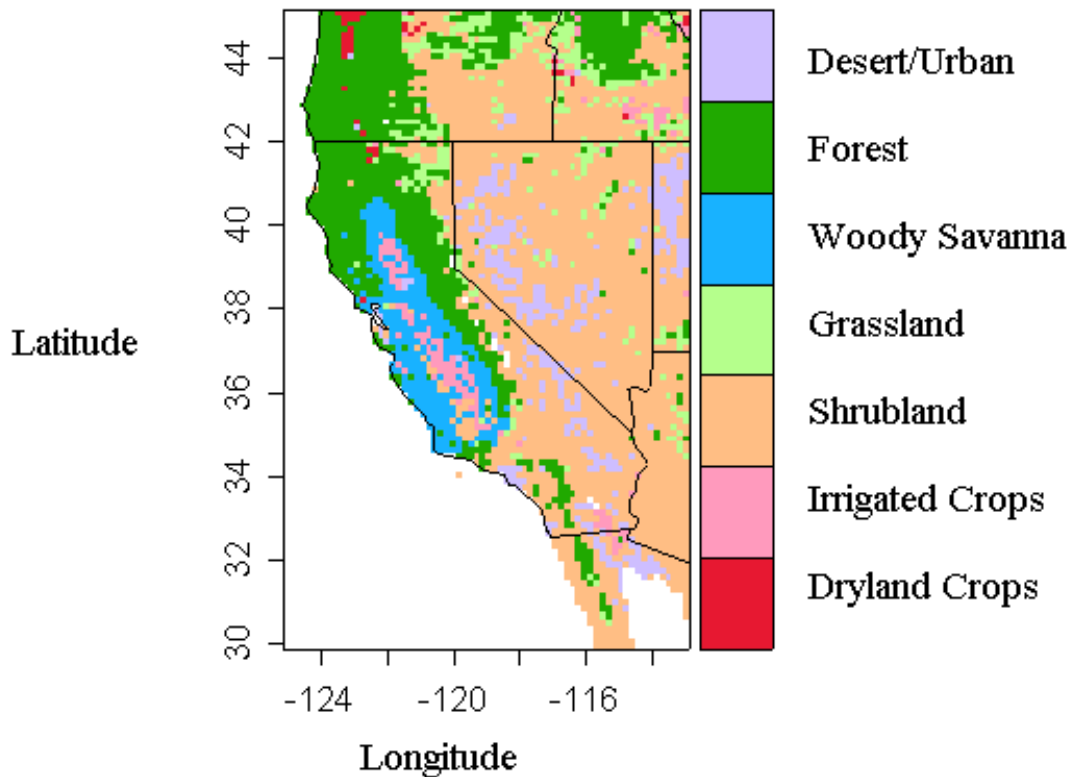


Figure 1. Vegetation distribution used to characterize land cover in the Western US for the simulations described in this report. The pixel size is 20 km on a side.

2.2 Prediction of Fossil Fuel CO₂ Emissions

We predicted the spatial and temporal distributions of CO₂ from fossil fuel emissions in the Western US by scaling estimated emissions of other combustion related pollutants (e.g., CO, and NO_x) that have been shown to correlate with atmospheric CO₂ concentration (Potosnak et al. 1999; Dillon et al. 2002; Gerbig et al. 2003; Lamarque et al. 2003). We used CO data reported in the US Environmental Protection Agency's (EPA) National Emission Inventory (NEI) of criteria

pollutants (EPA 2001). We obtained the NEI database from the Lake Michigan Air District Consortium (LADCO), on a grid with 36 km resolution for the continental US. The database contained hourly emissions for weekdays, Saturdays, and Sundays of each month in 1999. We extracted the CO data from the NEI and interpolated it to produce maps at 20 km resolution. Fossil fuel CO₂ emissions were calculated by multiplying CO emissions by a constant factor of 23.3, so that the annual sum of surface CO₂ emissions matched the Energy Commission's estimate of California CO₂ emissions in 1999 (363 Mt CO₂ minus approximately 17 Mt CO₂ for bunker and aviation fuel combustion) (Franco 2002). This scaling factor is reasonably consistent with previous work on atmospheric CO₂:CO ratios attributed to signals from fossil fuel combustion (Potosnak et al. 1999).

This method provides a high-resolution picture of fossil fuel CO₂ that is sufficient for this planning exercise. However, before using this type of information in an inverse analysis, refinements in the modeling should address the following issues. First, although the fossil CO₂ emission maps likely capture the dominant spatial and temporal features of the CO₂ emissions, they will contain regions and periods for which the CO₂:CO ratio varies (e.g., high CO from biomass burning or low CO from natural gas fired electricity production). A superior approach would be to use the information on fuel use underlying pollutant emissions to estimate CO₂ directly. Second, as mentioned above, these fossil fuel CO₂ emissions do not contain bunker or aviation CO₂ emissions from marine or aircraft transportation. These emissions can be included with some consideration of their spatial distribution in the atmosphere. Third, the predictions should use an emission inventory specific to the year being simulated or be adjusted appropriately.

2.3 Prediction of Concentration Signals

We predicted atmospheric CO₂ concentrations using the US EPA Community Multiscale Air Quality (CMAQ) model (<http://www.epa.gov/asmdnerl/models3/doc/science/science.html>). CMAQ was developed and applied for modeling air quality, incorporating the processes of bulk air motion, turbulent diffusion, and thermal and photo-chemistry (EPA 1999; Cohn et al. 2001; Bullock et al. 2002). Spatially and temporally resolved surface emissions were input to the model. Meteorological inputs were taken from the MM5 model predictions that included wind vectors, atmospheric boundary layer (ABL) height⁶, and vertical velocity fluctuations to compute the motion of air parcels.

CMAQ was used to calculate the temporal and spatial propagation of atmospheric CO₂ with the only sources and sinks occurring at the land surface as prescribed by the maps of NEE and fossil fuel CO₂ emissions. This is valid because CO₂ is not generated or lost through chemical reactions in the atmosphere. Initial and boundary conditions for CO₂ concentrations were set to a constant value of 370 ppm and signals were calculated with respect to this background. We subdivided NEE and fossil fuel CO₂ into in-state and out-of-state components and then calculated the CO₂ signals from four sources of surface flux: 1) NEE within CA (NEE.in), 2)

⁶ The height in the atmosphere to which air is well mixed.

NEE outside CA (NEE.out), 3) fossil fuel CO₂ from combustion inside CA (ffCO₂.in), and 4) fossil fuel CO₂ from combustion outside CA (ffCO₂.out).

2.4 Identification and Evaluation of Observation Stations

We identified an initial set of surface observation stations by starting with existing carbon cycle monitoring networks (e.g., NOAA-CMLD; <http://www.cmdl.noaa.gov>, and the Ameriflux Program <http://public.ornl.gov/ameriflux>). However, because of their relatively sparse coverage, we included two categories of other stations. Some of the additional stations were marine buoys planned by NOAA for offshore monitoring. Others were obtained from lists of radio and television transmission towers (<http://wireless2.fcc.gov/UlsApp/AsrSearch>) similar to the Park Falls Tower in Wisconsin where CO₂ measurements have been made previously (Bakwin et al. 1998).

We evaluated the efficacy of these observing stations by examining the time series concentration signals obtained at each station from the four component sources (NEE.in, NEE.out, ffCO₂.in, and ffCO₂.out). Three criteria were used in the selection process. First, we wanted to identify marine stations that had only small signals from the terrestrial sources. These stations would be used to quantify the CO₂ concentration boundary conditions for air entering the region. Second, we wanted to find stations that provided measurable signals from terrestrial sources. We used the time series data to calculate the mean mid-day signals from each component source and over each season. Mid-day CO₂ concentration signals were chosen because the mixing height of the atmospheric boundary layer is generally well defined and, although large, can be accurately represented in meteorological models. By comparison, nighttime atmospheric mixing is often poorly defined and represented. Third, we wanted the signals from NEE.in at the stations to be clearly distinguishable from signals due to the other sources. To quantify the degree of similarity between signals, we calculated the covariance between the signal from NEE.in. and the signals from each of the other three sources. If there was little covariance, then signal variations due to NEE.in could be measured and the signals from other sources would add noise but not bias to the measurement.

3. Results and Discussion

3.1 Spatiotemporal distribution of Net Ecosystem Exchange

Using the coupled MM5-LSM1 simulations, we predicted regional meteorology and NEE for the Western region over the first 15 days of March, June, September, and December. We calculated the average NEE for each cover type over the four periods of the study. Both the mean NEE for forest (-0.66 t C ha⁻¹ yr⁻¹) and for shrub lands (0.38 t C ha⁻¹ yr⁻¹) were close to carbon balance, and reasonably consistent with the range of estimates for the annual NEE from the literature (Goldstein et al. 2000; Baldocchi et al. 2001; Falge et al. 2002; Law et al. 2002). The mean NEE of irrigated crops (-10.5 t C ha⁻¹ yr⁻¹) and dry crops (-5.5 t C ha⁻¹ yr⁻¹) both showed the large net uptake expected from managed agricultural systems, where nitrogen is applied and crop residues are removed after harvest. The predicted mean NEE of woody savanna (-19.5 t C ha⁻¹ yr⁻¹) and grassland (-6.2 t C ha⁻¹ yr⁻¹) were sufficiently different from values observed in California oak woodland (pers. comm. D. Baldocchi) and grassland systems (Xu et al. 2004) that adjustments

were made to reduce net daytime uptake and increase net nighttime release as described in section 2.1. The resulting mean NEE of woody savanna and grassland were both $-0.7 \text{ t C ha}^{-1} \text{ yr}^{-1}$.

Predicted NEE exhibits large variations both in time and with cover type. For example, the maps in figure 2 (and figure A1, appendix A) show the diurnal variation in NEE over the course of a day in June. As expected, nighttime NEE is positive, reflecting plant and soil respiration of CO_2 to the atmosphere. Daytime NEE is negative for ecosystems that are actively assimilating CO_2 from the atmosphere (e.g., forest, woodland, and crop lands), but slightly positive for shrub ecosystems where, on this day, respiration was larger than photosynthesis. The large spatial variations in NEE reflect the varied ecosystems present in California. Similarly, the ecosystems have different seasonality, which generates seasonal NEE patterns (see figure A2 in Appendix A). Capturing these seasonal NEE variations is important in order to accurately predict seasonal variations in atmospheric CO_2 concentrations.

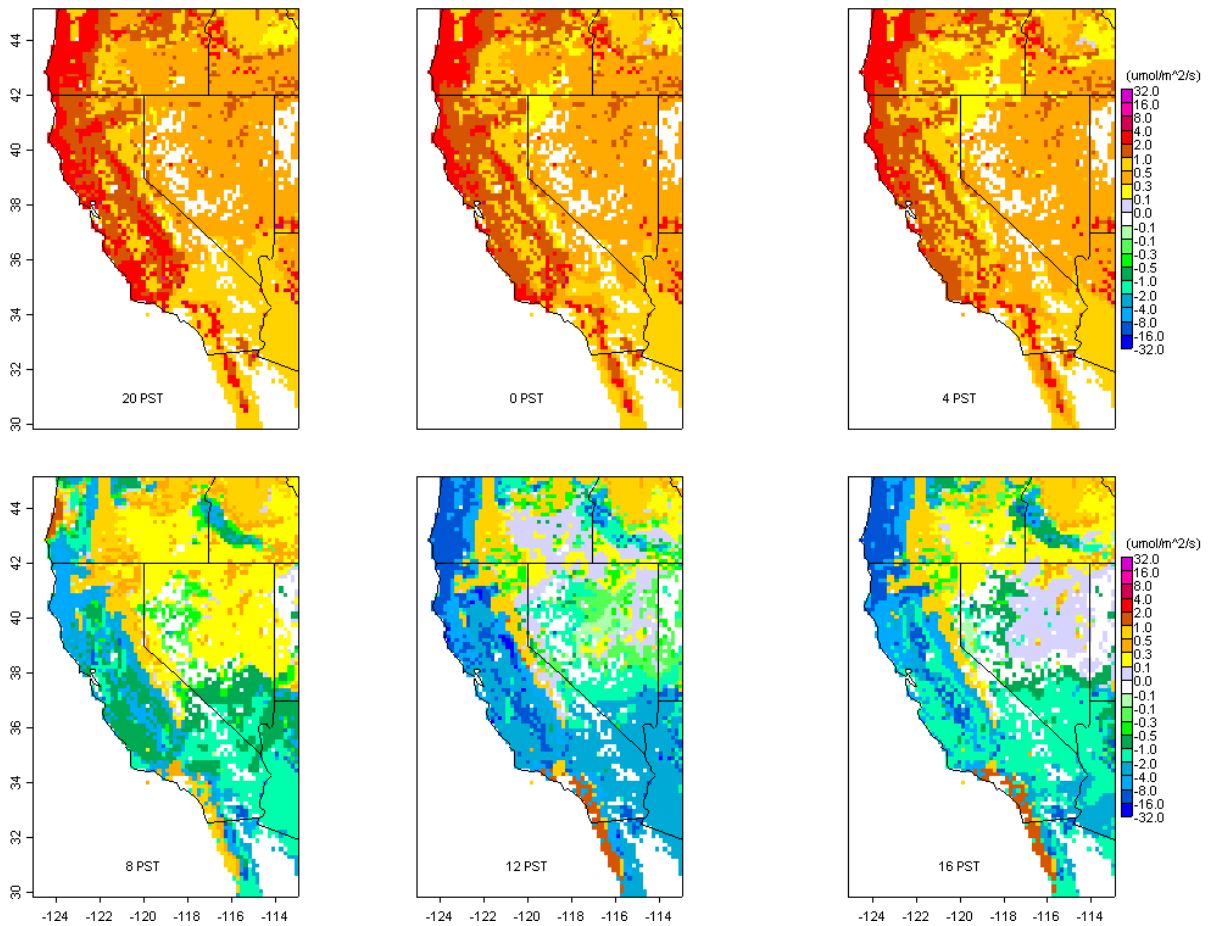


Figure 2. Maps showing the diurnal cycle of predicted net ecosystem exchange, NEE ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$), for a day in June. Horizontal and vertical axes are longitude and latitude respectively. Upper panels show positive fluxes to the atmosphere dominate at night due to respiration. Lower

panels show negative fluxes to the atmosphere dominate during the day when photosynthesis removes CO₂ from the atmosphere. Note: NEE is shown with a logarithmic scale.

3.2 Spatiotemporal distribution of fossil fuel CO₂ emissions

A representative diurnal cycle of CO₂ emissions from fossil fuel combustion in the Western US is shown in Figure 3 for a weekday in June. As expected, fossil fuel CO₂ emissions were strong in large urban centers (larger than any positive ecosystem fluxes), but were spatially localized. Away from urban centers, emissions remained significant only along suburban transportation corridors. Emissions from aircraft and ship traffic were not included, but represent a small fraction of estimated CO₂ emissions from California. The temporal variations in fossil fuel CO₂ fluxes were much smaller than temporal variations in signals from NEE because ecosystems have distinct diurnal and seasonal cycles of net uptake and release. Given that the temporal variations in CO₂ concentration signals from fossil fuel CO₂ are smaller than the variations from NEE we hypothesize that fossil fuel signals could be predicted and subtracted from measured signals to yield the signal from NEE. This hypothesis should be investigated in future work.

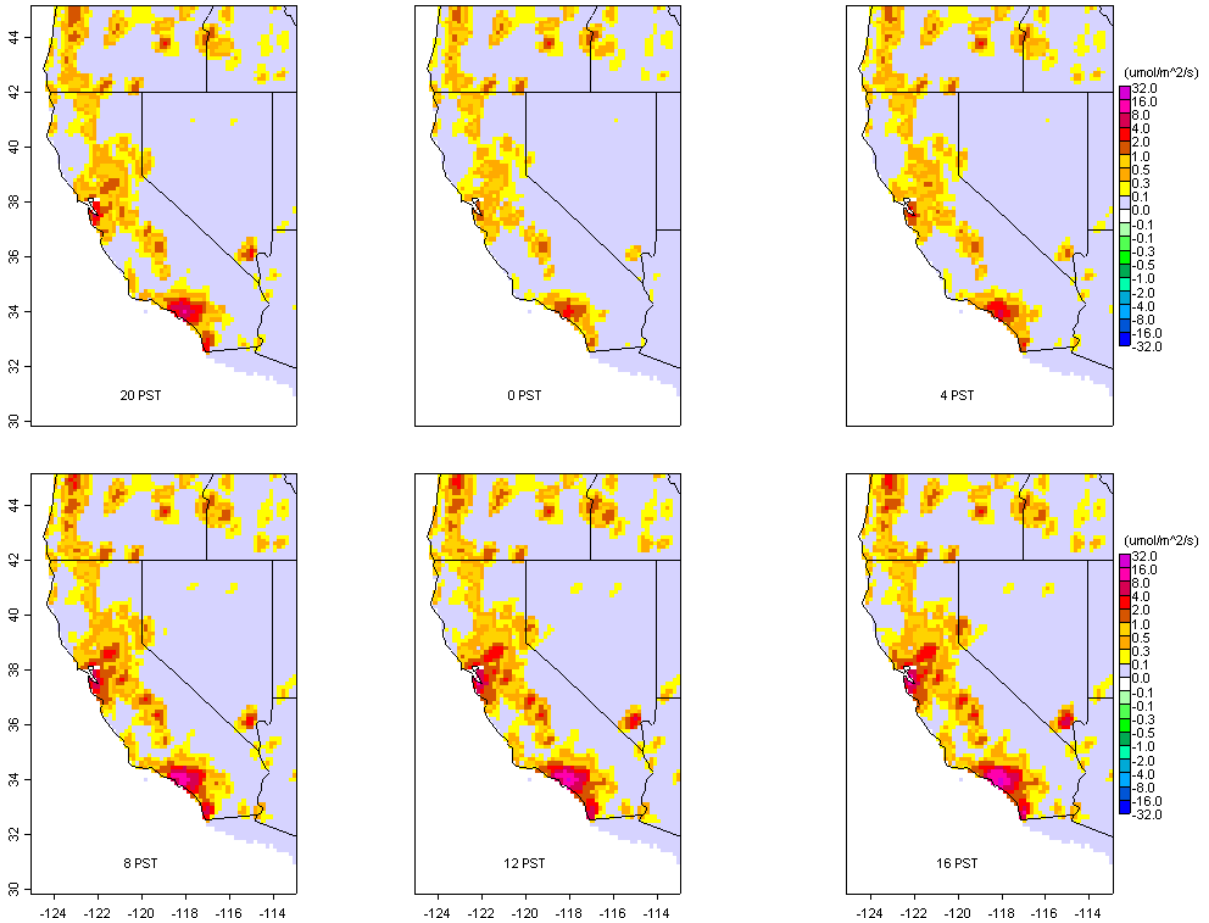


Figure 3. Maps showing the diurnal cycle of estimated fossil fuel CO₂ emissions for the Western US for a weekday in June. Panel layout, axes, and logarithmic scale for units are the same as in Figure 2.

3.3 Spatiotemporal distribution of CO₂ concentrations in the atmosphere

Atmospheric CO₂ concentrations were predicted for the four simulation periods using predicted NEE, fossil fuel CO₂ emissions, and meteorology described above. All CO₂ concentrations are reported as deviations from the constant boundary condition applied at the edge of the simulation domain. Figure 4 shows the diurnal cycle in CO₂ concentration signals from NEE in the lowest (0-100 m) model layer of the atmosphere for the day shown in Figure 2. In areas with strong surface uptake, concentrations were reduced below the background, while in areas with strong sources they increased above the background. Air parcels containing significant CO₂ concentration signals were predicted to advect to marine locations up to 300 km west of the coast, consistent with the results of another study (Riley et al. 2004). This suggests that attempts to use the data from marine and coastal stations to provide boundary condition information on the air entering California will require correction for terrestrial influence. From these maps we can also conclude that the spatial differences in CO₂ signals across CA are large enough to be clearly measured with available instrumentation.

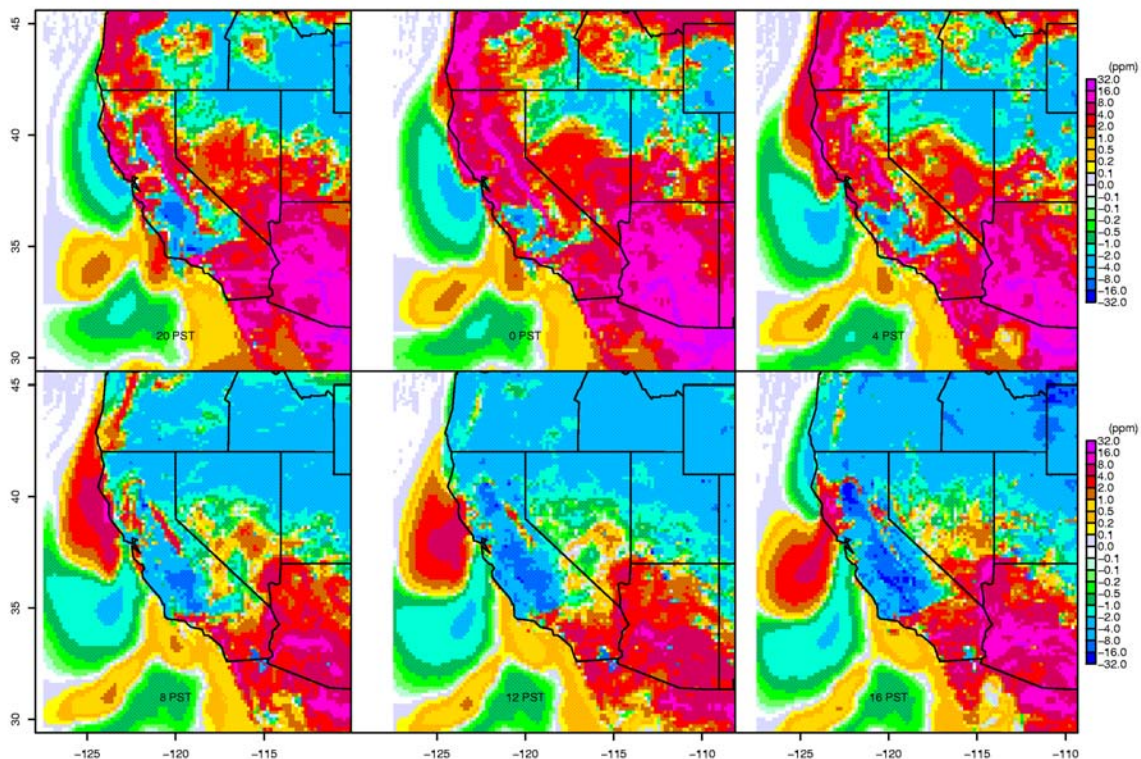


Figure 4. Maps showing diurnal variation in surface layer CO₂ concentrations resulting from NEE for the day shown in Fig 2. The upper (lower) three panels correspond to nighttime

(daytime) periods. Horizontal and vertical axes are longitude and latitude respectively. Note that CO₂ concentration contours are logarithmically scaled. The spatial variations in CO₂ concentration are large enough to be clearly detected with available instrumentation.

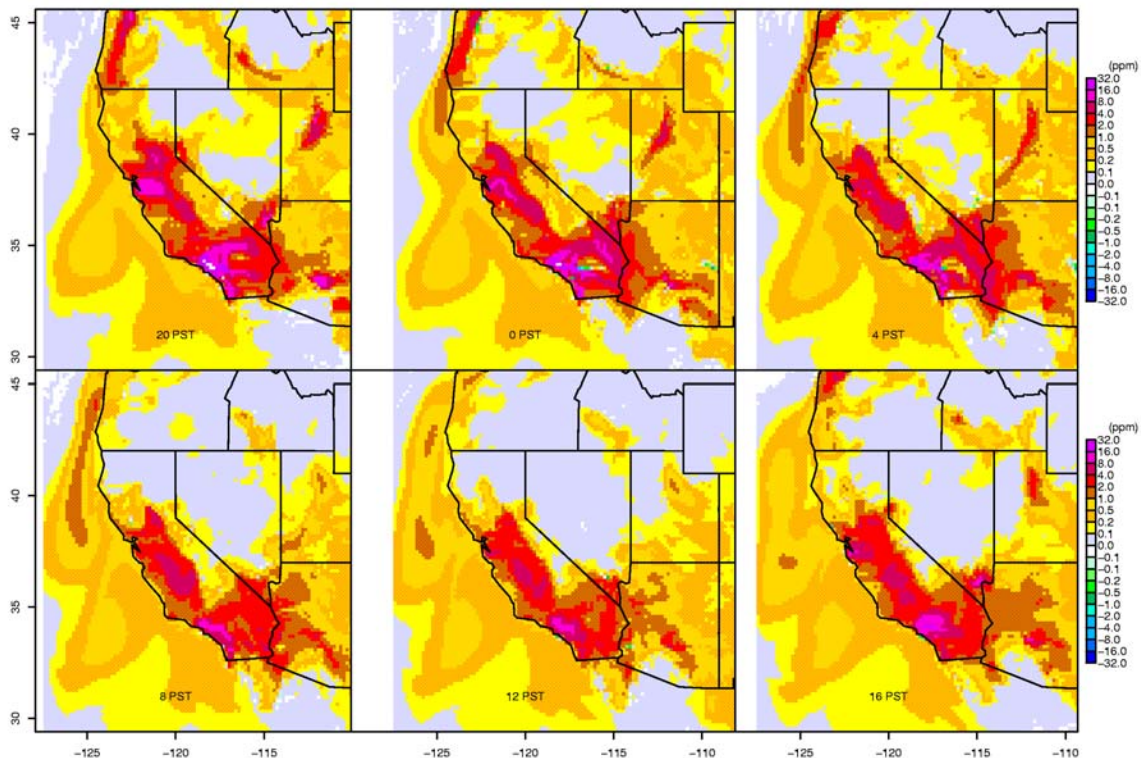


Figure 5. Maps of diurnal variation in surface layer CO₂ concentrations resulting from fossil fuel CO₂ emission for June time period as in Figure 4. Panel layout, axes, and logarithmic scale for units are the same as in Figure 4.

Figure 5 shows the diurnal cycle in CO₂ concentrations resulting from fossil fuel CO₂ emissions for the same period as Figure 4. In this case, only increases above background are predicted because all fluxes are positive with respect to the atmosphere. As with the emissions map, the most significant increments in CO₂ concentration occurred near the urban centers, but significant plumes of CO₂ enriched air were also transported to other locations far from the location of emissions. However, the mean and diurnal variations in fossil CO₂ signals were generally smaller than CO₂ signals from NEE at most locations away from urban centers. The smaller size and lower variability of fossil CO₂ signals suggests that subtracting predicted fossil CO₂ signals from future measurements of CO₂ concentration may be effective for estimating the signals from NEE at many locations. This increases our confidence that atmospheric observations will enable effective estimation of NEE.

In order to facilitate identification of the most promising locations for observing stations, we calculated the CO₂ signals for four components: signals derived from 1) NEE within CA

(NEE.in), 2) NEE outside CA (NEE.out), 3) fossil fuel CO₂ from combustion inside CA (ffCO₂.in), and 4) fossil fuel CO₂ from outside CA (ffCO₂.out). Figure 6 shows the mean midday surface layer concentration signals for each of the four components in June. The panels of Figure 6 show that the central valley and central coast areas had negative signals from NEE inside CA and positive signals from fossil fuel CO₂ produced inside CA. During this period Southern CA had positive signals from both NEE and fossil fuel CO₂. Coastal and marine areas obtained smaller positive signals from plumes transported off the coast. As shown in Appendix B, these patterns varied significantly with season. In the case of NEE, the patterns varied because NEE has significant seasonal variations due to ecology and climate. In the case of fossil fuel CO₂, the variations are driven mostly by variations in the mean meteorology for the different months.

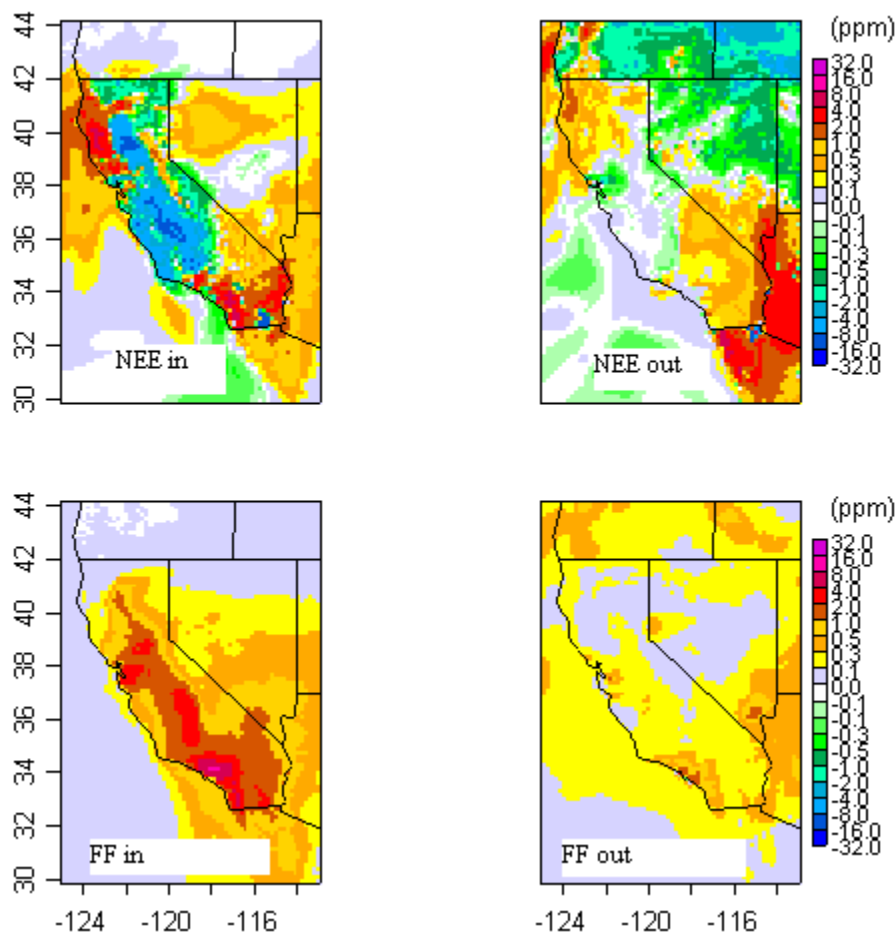


Figure 6. Maps of June midday mean surface layer CO₂ concentrations resulting from NEE inside (upper left) and outside (upper right) California, and fossil CO₂ emissions inside (lower left) and outside (lower right) California. As with previous figures the horizontal and vertical axes are longitude and latitude respectively. Note that CO₂ concentration contours are logarithmically scaled.

3.4 Identification and Evaluation of Observation Stations

3.4.1 Existing and Potential Future Stations

A set of existing and potential future monitoring stations is shown in Figure 7 and Table 1. Some of the existing stations are part of the NOAA CMDL measurement and sampling network (see the following report: <http://www.cmdl.noaa.gov/publications/annrpt27/>), some are independent sites (e.g., Scripps), and others are Ameriflux sites where accurate CO₂ measurements could be implemented, and have been at Blodgett Forest (station 13). We note that there are other existing ecological and biogeochemical flux measurement sites in California that could be included in future analysis, but such a compilation was beyond the scope of this study. The marine stations are chosen to sample the atmosphere from relatively pristine locations that provide boundary conditions for air entering the California. The coastal stations sample air in a transition regime. The inland stations sample air that is heavily influenced by terrestrial surface fluxes. We did not include a significant number of sites East of California because the topography of the Sierra poses a serious challenge for modeling atmospheric transport.

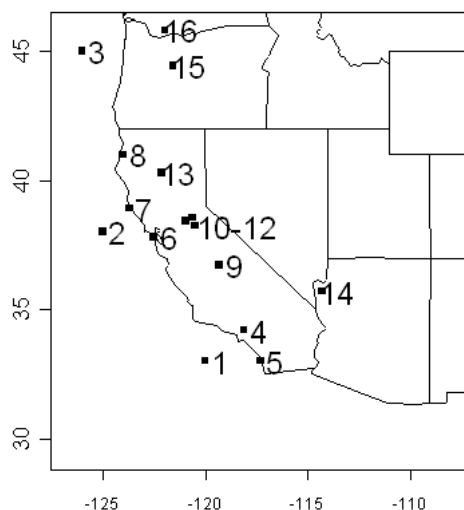


Figure 7. The positions of existing and possible future observations stations examined in this study. The locations range from marine stations that largely sample background air to inland stations that sample air subject to strong terrestrial influence.

Table 1. Existing and potential monitoring stations. For existing sites the name of the research program is listed in parentheses.

Station	Name	Lat	lon	agl (m)
1	Buoy	33.000	-120.000	3
2	Buoy	38.000	-125.000	3
3	Buoy	45.000	-126.000	3
4	Tall Tower	34.200	-118.100	296

5	Scripps Pier (Scripps Inst.)	33.000	-117.300	10
6	Tall Tower	37.800	-122.500	298
7	Point Arena (NOAA)	38.911	-123.692	10
8	Trinidad Head (NOAA)	41.000	-124.000	10
9	Tall Tower	36.700	-119.300	259
10	Tall Tower	38.240	-120.502	624
11	Tonzi Ranch (Ameriflux)	38.432	-120.966	12
12	Blodgett Forest (Ameriflux)	38.550	-120.633	15
13	Tall Tower	40.300	-122.100	304
14	Tall Tower	35.700	-114.300	396
15	Metolious (Ameriflux)	44.437	-121.567	12
16	Wind River (Ameriflux)	45.821	-121.951	73

3.4.2 CO₂ concentration signals at monitoring station locations

The time series of concentration signals calculated for each station was extracted from the concentration maps and examined to determine the contributions to the predicted signals from NEE and fossil CO₂ emissions inside and outside of California. Figure 8 (and Appendix C) shows the mean midday concentration signal for each station and season. As mentioned above, we separated the stations into categories based on whether their primary influence was oceanic (stations 1-3) and primarily of use for monitoring boundary conditions, coastal (4-8) and influenced by a mixture of marine and interior air, interior (9-13) and strongly influenced by NEE and fossil CO₂, or outside of California (14-16) and of value for monitoring boundary conditions. Within the station groupings, the stations are arranged from South to North.

The general patterns show increasing signals from NEE.in and decreasing signals from ffCO₂.in from south to north. This is because Southern CA is quite arid and does not support the high rates of ecosystem carbon exchange that are possible in the comparatively moister conditions found in Northern CA. Other significant features for stations within CA are consistent seasonality (also see Appendix C) with negative NEE.in signals during March, small or positive signals in June near the coast but negative signals inland (from irrigated agriculture and Sierran forests), uniformly positive signals in September when most areas are subject to water stress, and small signals in December when temperatures are lower. Patterns for ffCO₂.in and out show less seasonality, largely due to variations in atmospheric transport because fossil emissions did not vary significantly with season.

The signals predicted for the different stations vary by grouping and latitude. Stations 1 and 2 show mean signals of ~ 1-2 ppm with comparable day-to-day variations, suggesting that their use as background monitoring stations will require incorporating estimated terrestrial signals. We note that this may also be true for station 3; it is near the northwest edge of the modeling domain and signals from terrestrial CO₂ exchanges even farther north may influence the signal at station 3. Stations 6-13 showed significant signals from NEE.in, while stations 4, 5, 9, 11, and 12 all exhibited large signals from ffCO₂.in and could be used to monitor fossil emissions. Stations 14-16 exhibited strong signals from NEE.out and ffCO₂.out and would therefore provide information to estimate a combination of those sources. Finally, the signals from the tall towers

(4, 6, 9, 10, 13, 14) generally exhibited smaller day-to-day variations as compared with nearby surface stations (Appendix B). This reduced variability occurs because they average over larger areas, a fact that makes them potentially superior for regional inverse studies.

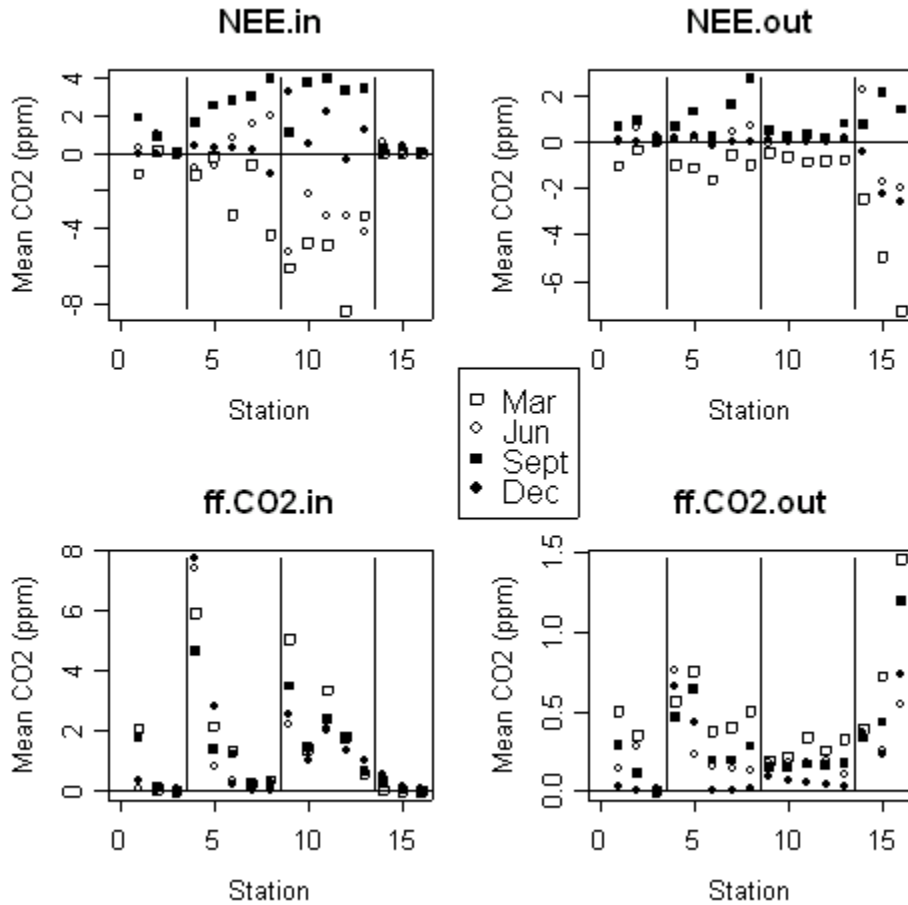


Figure 8. Mean CO₂ concentration signals from NEE in and out and fossil fuel CO₂ emissions in and out of California for each of the four months examined in this study. Vertical lines separate the stations into four groups in the following order: oceanic, coastal, interior, and interior outside California. CO₂ concentrations are deviations from the seasonal background.

3.4.3 How well do concentration signals at monitoring sites allow discrimination of NEE inside CA from other signals?

Additional insight into the relative strengths of the observing stations can be obtained by inspecting the covariance between signals from NEE.in and the other signals shown in Figure 9. We normalized the covariance between signals from NEE.in and other signals by the standard deviation in signal from NEE.in. This measure shows the amount of signal variation from the other sources that are correlated with (and could not be distinguished from) variations from

NEE.in. In most cases, the covarying signals were small compared to the signal variations from NEE.in. This suggests that if the variables governing variation in the signal from NEE.in can be captured in an ecosystem-atmosphere model (e.g., MM5-LSM1), then an inverse model can use the measured signals to estimate NEE.in without interference from the other signals. A subset of stations (6, 7, 8, 10, 12, and 13) did not exhibit significant signal correlations between NEE.in and ffCO₂.in, and only limited correlations with NEE.out and ffCO₂.out. Station 12 (Blodgett forest) appeared particularly well suited in this respect, although modeling the transport fields for Blodgett poses a greater challenge (because of the mountainous terrain) than for the Central Valley sites. Based on these results, we conclude that measurements made at a subset of these stations are likely to offer an opportunity to quantify NEE.in without significant interference from other sources.

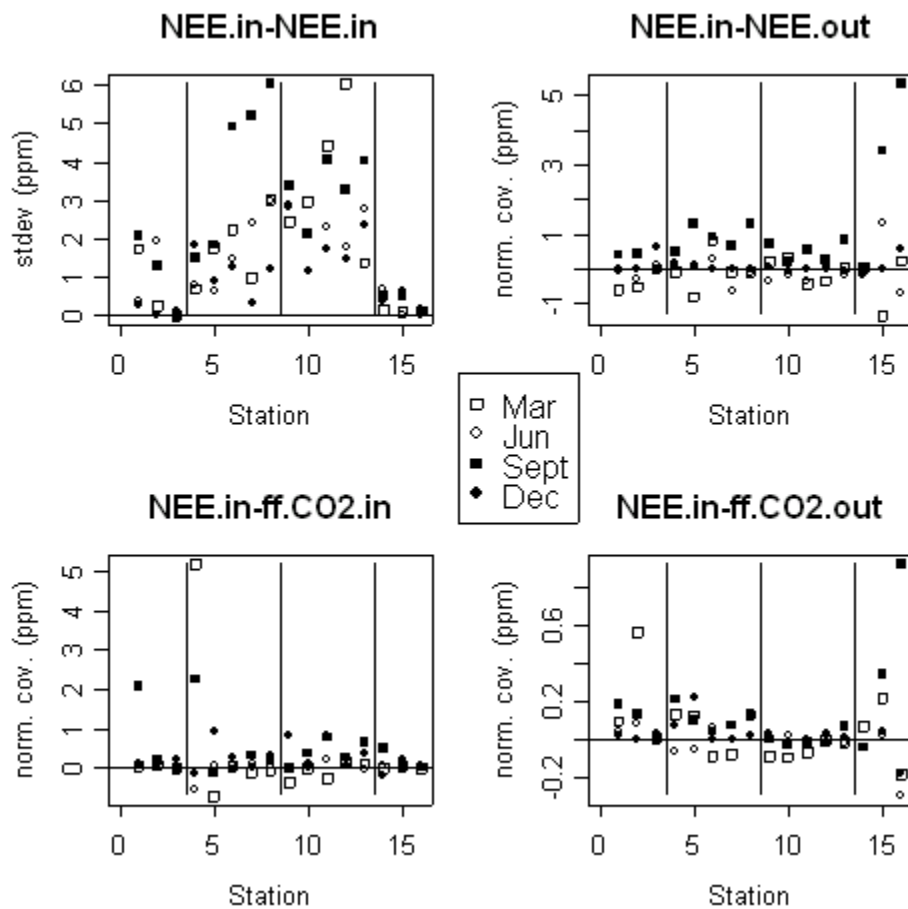


Figure 9. Standard deviation of midday CO₂ signals from NEE inside California (top left), showing predicted variation in daytime signals for each season. The normalized covariance of the signal from NEE.in with the other signals is also shown: NEE outside CA (top right), ffCO₂ inside CA (bottom left), and ffCO₂ outside CA (bottom right). In all cases the normalization factor is the standard deviation of NEE.in, yielding the component (or projection) of each signal that is correlated with signal from NEE.in. Several stations show normalized signal covariances

that are small compared to the standard deviation of the NEE.in signal are hence likely to offer an opportunity to quantify NEE.in without interference from other sources.

4. Conclusions and Recommendations

This section provides conclusions and recommendations for 1) a proposed set of initial observation stations, sampling frequencies, and a measurement protocol for CO₂, 2) analysis of the uncertainties in future inverse model estimates of NEE and refinement of observing stations, 3) measurement of additional trace gases that would enhance our ability to estimate CO₂ exchange, 4) initiation of a data synthesis effort to couple data from different sources relevant for estimating NEE, and 5) application of atmospheric methods for non-CO₂ GHG emissions.

4.1 Observation Stations and Sampling Frequency

The results of this project have shown that there are locations where atmospheric CO₂ measurements will help constrain California's NEE. Based on these results, we recommend an initial sampling at a limited number of sites to provide a baseline for future efforts:

- New terrestrial stations: New observing stations should be deployed at existing communications tall towers in California, chosen so that CO₂ signals are strongly influenced by either NEE.in or ffCO₂.in. Stations 4, 9, 10, and 13 are well suited in this respect. Station 4 would provide an unambiguous measure of emissions from the Los Angeles area; station 9 would sample air in the San Joaquin Valley that is effected by a mixture of agricultural and Bay Area emissions ; station 10 would be effective for monitoring signals from NEE due to Central Valley agriculture and forests in the Sierra, with a comparatively small signal from fossil CO₂ emissions; station 13 would monitor exchange from northern valley agriculture and forests.
- Sampling frequency: Ideally air sampling should be conducted continuously in an automated manner to accurately capture the concentration during the well-mixed afternoon period and the diurnal variation. Data should be collected and averaged into approximately 30-minute intervals for further analysis. However, if resources are limited, periodic sampling with NOAA flask systems could be performed to provide a valuable initial data to test the model predictions similar to those performed in this study.
- Measurement protocol: Air sampling, instrument design, calibration, and analysis should follow established methods to ensure accuracy and minimize systematic errors (Bakwin et al. 1995). In particular, continuous measurements should be calibrated against multiple known standards (preferably supplied by NOAA) several times per day. Periodic sampling with NOAA flask systems should be performed at these sites to provide a check on the continuous measurements.
- Additional stations: Observations are planned for a variety of platforms to monitor background air entering California for carbon cycle and air quality purposes. This will likely include the three planned marine stations planned by NOAA, periodic aircraft profiling planned by NOAA for Trinidad Head and a site in Southern CA, and potentially

continuous or flask sampling from one or more mountain-top stations. As noted above, interpreting signals from the marine stations will require estimation of signals from terrestrial sources in North America. It is also expected that existing flux measurement sites will be augmented with accurate and precise CO₂ concentration measurements. These data sets should be included in future analyses.

4.2 Uncertainty Analysis and Site Optimization

Because it was beyond the scope of the current work, we did not include an analysis of the uncertainty in NEE that could be obtained using data from the suggested observing stations in an inverse approach. Before implementing a full set of observing stations, an analysis of the likely uncertainty in estimating NEE using atmospheric concentration data is warranted. In particular, one should determine whether the uncertainty in estimating NEE using the atmospheric inverse approach is very large or an infeasible number of stations are required. This is a complicated problem because the choice of the inverse approach and multiple sources of uncertainty all affect the final uncertainty in CO₂ flux as well as the optimal locations for a given number of stations. In particular, the desired spatial resolution of the NEE estimates will depend on the number and locations of the stations. Sources of (likely correlated) uncertainty include: 1) the *a priori* NEE estimate, 2) the meteorology and hence atmospheric transport, 3) the background CO₂ concentrations from outside the modeling domain, 4) the CO₂ contribution from fossil emissions, and 5) the uncertainty in the concentration measurements themselves. To address these issues we recommend:

- First, simplify the consideration to focus on the uncertainty in estimating the total NEE from inside California. Use simplified forms (assume errors are uncorrelated) for the uncertainties from each of the five above sources, apply a Monte Carlo approach to generate pseudo data, and then estimate the distribution of California NEE obtained from applying an inverse method. Iterate this calculation with different combinations of towers or alternate locations to estimate the optimal number and location of towers, in part driven by budget considerations. This approach would probably provide a good first guess as to whether acceptable uncertainties in NEE for California could be obtained with a limited number of towers and hence whether and where to proceed with instrument deployment.
- Second, improve the understanding of each of the five above sources of uncertainty. For example, predictions of CO₂ concentrations should include calculation of signals from the seasonally and spatially varying global background of CO₂ concentrations. First order estimates in boundary conditions could be derived from the CMDL Globalview CO₂ product (<http://www.cmdl.noaa.gov/ccgg/globalview/index.html>). These estimates should be compared with the results of aircraft campaigns that measure the vertical profile of CO₂ (and other gases) over the Eastern Pacific. Estimating the boundary conditions will also be valuable because it will allow a comparison between predicted CO₂ concentrations and those measured at existing stations.
- Third, expand the time periods covered in the investigation and the spatial resolution of the inverse estimates. Include several complete annual cycles to sample different weather patterns over multiple years which capture inter-annual variations in both meteorology and

NEE (which is strongly dependent on meteorological forcing). Investigate how well NEE from different regions within California can be separately identified.

4.3 Additional Trace Gas Species for CO₂ Source Attribution

Although it is beyond the scope of this study to analyze this problem in detail, there are several additional trace-gas species that would provide important additional constraints on CO₂ exchange:

- CO and certain organic compounds (e.g., C₂H₂) have been identified as useful tracers of combustion (Potosnak et al. 1999) and should be developed for use in California, as is currently the case at Blodgett Forest. The use of CO (and other pollutant) measurements from the California Air Resources Board (CARB) might play a valuable part in this effort, particularly near urban areas where pollutant concentrations are high enough to be clearly detected with the available instruments. If dedicated future CO₂ measurement stations are commissioned, inclusion of precise monitoring CO (and other species) in addition to CO₂ would significantly enhance their capability.
- Radiocarbon CO₂ (¹⁴CO₂) can provide a strong constraint on fossil fuel combustion because the 5700 year half-life for ¹⁴C is short compared to the age of geologically derived petroleum hydrocarbons and fossil hydrocarbons contain no ¹⁴C. Given the current precision and accuracy of ¹⁴CO₂ measurements, it should be possible to quantify the ¹⁴CO₂ free component of air to ~ 1 ppm. NOAA CMDL and LBNL are independently initiating the use of atmospheric sampling and analysis of ¹⁴CO₂ for this purpose. Because of the high cost for precision analysis (roughly \$300 per sample) this approach is currently beyond the scope of most routine monitoring programs, but might be employed to add precision to fossil CO₂ estimates during an intensive measurement period.
- The ratio of CO₂ to ²²²Rn has been used to estimate net regional CO₂ exchange using the assumption that surface emission rate of ²²²Rn can be accurately determined (Biraud et al. 2000; Biraud et al. 2002). This technique is well suited to the problem of CA carbon budgets because much of the air entering CA is of oceanic origin and will be depleted in ²²²Rn. Application of this technique will require detailed analysis of ²²²Rn source fields for CA because work has shown that ²²²Rn emissions from soil can vary substantially with meteorology and soil moisture (Conen et al. 2002; Levin et al. 2002).

4.4 Synthesis of Bottom-Up and Top-Down Methods

A crucial aspect of using atmospheric, or “top-down”, methods to monitor California’s CO₂ emissions will be to sensibly incorporate “bottom-up” information (e.g., ecosystem production from forest inventories and crop yields, fossil CO₂ emissions from improved inventory methods, eddy covariance CO₂ flux data from towers and aircraft flights, and ecosystem models of CO₂ exchange) in the inversion. This approach is an emerging area and developments in the NACP will be valuable to addressing these issues in CA.

4.5 Potential for Monitoring Non-CO₂ GHG Emissions

The approach that we have applied here could also be profitably applied to quantify the atmospheric signals generated by surface emission of non-CO₂ GHGs. Because the relative uncertainties for fluxes of non-CO₂ GHGs are considerably larger than for CO₂ (Franco 2002), atmospheric budget (Biraud et al. 2002) and inverse approaches may be particularly effective in significantly reducing those uncertainties using atmospheric observations.

References

- Bakwin, P. S., K. J. Davis, C. Yi, S. C. Wofsey, J. W. Munger, L. Haszpra and Z. Barcza (2004). Regional carbon dioxide fluxes from mixing ratio data. Tellus B **56**(4): 301-311.
- Bakwin, P. S., P. P. Tans, D. F. Hurst and C. L. Zhao (1998). Measurements of carbon dioxide on very tall towers: results of the NOAA/CMDL program. Tellus Series B-Chemical & Physical Meteorology **50**(5): 401-415.
- Bakwin, P. S., P. P. Tans, C. L. Zhao, W. Ussler and E. Quesnell (1995). Measurements of Carbon Dioxide on a Very Tall Tower. Tellus Series B-Chemical & Physical Meteorology **47**(5): 535-549.
- Baldocchi, D., et al. (2001). FLUXNET: A new tool to study the temporal and spatial variability of ecosystem-scale carbon dioxide, water vapor, and energy flux densities [Review]. Bulletin of the American Meteorological Society **82**(11): 2415-2434.
- Betts, A. K. and J. H. Ball (1998). Fife Surface Climate and Site-Average Dataset 1987-89. Journal of the Atmospheric Sciences **55**(7): 1091-1108.
- Biraud, S., et al. (2002). Quantification of carbon dioxide, methane, nitrous oxide and chloroform emissions over Ireland from atmospheric observations at Mace Head. Tellus Series B-Chemical and Physical Meteorology **54**(1): 41-60.
- Biraud, S., et al. (2000). European greenhouse gas emissions estimated from continuous atmospheric measurements and radon 222 at Mace Head, Ireland. Journal of Geophysical Research-Atmospheres **105**(D1): 1351-1366.
- Bonan, G. B. (1995). Land-atmosphere CO₂ exchange simulated by a land surface process model coupled to an atmospheric general circulation model. Journal of Geophysical Research (Atmospheres) **100**(D2): 2817-2831.
- Bonan, G. B. (1996). A land surface model (LSM version 1.0) for ecological, hydrological, and atmospheric studies: Technical description and user's guide. Boulder, CO, NCAR: 150.
- Bonan, G. B., K. J. Davis, D. Baldocchi, D. Fitzgerald and H. Neumann (1997). Comparison of the NCAR LSM I land surface model with BOREAS aspen and jack pine tower fluxes. Journal of Geophysical Research **102**(C12): 29,065-29,076.
- Bullock, O. R. and K. A. Brehme (2002). Atmospheric mercury simulation using the CMAQ model: formulation description and analysis of wet deposition results. Atmospheric Environment **36**(13): 2135-2146.
- Chen, F. and J. Dudhia (2001). Coupling an advanced land surface-hydrology model with the Penn State-NCAR MM5 modeling system. Part II: Preliminary model validation. Monthly Weather Review **129**(4): 587-604.
- Ciais, P., P. P. Tans, M. Trolier, J. W. C. White and R. J. Francey (1995). A Large Northern Hemisphere Terrestrial Co₂ Sink Indicated by the C-13/C-12 Ratio of Atmospheric Co₂. Science **269**(5227): 1098-1102.
- Cohn, R. D., B. K. Eder, S. K. Leduc and R. L. Dennis (2001). Development of an aggregation and episode selection scheme to support the Models-3 Community Multiscale Air Quality model. Journal of Applied Meteorology **40**(2): 210-228.
- Conen, F. and L. B. ROBERTSON (2002). Latitudinal distribution of radon-222 flux from continents. Tellus B **54**(2): 127-133.
- Cooley, H. S., W. J. Riley, M. S. Torn and Y. He (2004). Impact of agricultural practice on regional climate in a coupled land surface mesoscale model. in Press J. Geophys. Res.

- Crisp, D., et al. (2004). The orbiting carbon observatory (OCO) mission. Trace Constituents in the Troposphere and Lower Stratosphere. **34**: 700-709.
- Denning, A. S., G. J. Collatz, C. G. Zhang, D. A. Randall, J. A. Berry, P. J. Sellers, G. D. Colello and D. A. Dazlich (1996). Simulations of Terrestrial Carbon Metabolism and Atmospheric Co₂ in a General Circulation Model .1. Surface Carbon Fluxes. Tellus Series B-Chemical & Physical Meteorology **48**(4): 521-542.
- Denning, A. S., D. A. Randall, G. J. Collatz and P. J. Sellers (1996). Simulations of Terrestrial Carbon Metabolism and Atmospheric Co₂ in a General Circulation Model .2. Simulated Co₂ Concentrations. Tellus Series B-Chemical & Physical Meteorology **48**(4): 543-567.
- Desjardins, R. L., et al. (1997). Scaling up Flux Measurements for the Boreal Forest Using Aircraft-Tower Combinations. Journal of Geophysical Research-Atmospheres **102**(D24): 29125-29133.
- Dillon, M. B., M. S. Lamanna, G. W. Schade, A. H. Goldstein and R. C. Cohen (2002). Chemical evolution of the Sacramento urban plume: Transport and oxidation - art. no. 4045. Journal of Geophysical Research-Atmospheres **107**(D5-D6): 4045.
- EPA (1999). Science Algorithms of the EPA Models-3 Community Multiscale Air Quality (CMAQ) Model, EPA.
- EPA (2001). Inventory of U.S. Greenhouse Gas Emissions and Sinks: 1990-1999. Washington, DC,, U.S. Environmental Protection Agency.
- Falge, E., et al. (2002). Phase and amplitude of ecosystem carbon release and uptake potentials as derived from FLUXNET measurements. Agricultural and Forest Meteorology **113**(1-4): 75-95.
- Franco, G. (2002). Inventory of California Greenhouse Gas Emissions and Sinks: 1990-1999, California Energy Commission.
- Gerbig, C., J. C. Lin, S. C. Wofsy, B. C. Daube, A. E. Andrews, B. B. Stephens, P. S. Bakwin and C. A. Grainger (2003). Toward constraining regional-scale fluxes of CO₂ with atmospheric observations over a continent: 2. Analysis of COBRA data using a receptor-oriented framework - art. no. 4757. Journal of Geophysical Research-Atmospheres **108**(D24): 4757.
- Goldstein, A. H., et al. (2000). Effects of climate variability on the carbon dioxide, water, and sensible heat fluxes above a ponderosa pine plantation in the Sierra Nevada (CA). Agricultural & Forest Meteorology **101**(2-3): 113-129.
- Grell, G., J. Dudhia and D. Stauffer (1995). A description of the fifth-generation Penn State/NCAR mesoscale model (MM5). Boulder, CO, NCAR.
- Gurney, K. R., et al. (2002). Towards robust regional estimates of CO₂ sources and sinks using atmospheric transport models. Nature **415**(6872): 626-630.
- Hayhoe, K., et al. (2004). Emissions pathways, climate change, and impacts on California. Proceedings of the National Academy of Sciences of the United States of America **101**(34): 12422-12427.
- Helliker, B. R., et al. (2004). Estimates of net CO₂ flux by application of equilibrium boundary layer concepts to CO₂ and water vapor measurements from a tall tower. JGR-Atmospheres **109**: D20106.
- IPCC (2001). 2001: Observed Climate Variability and Change. In: Climate Change 2001: The Scientific Basis. Contribution of Working Group I to the Third Assessment Report of the

- Intergovernmental Panel on Climate Change. New York, NY, USA, Cambridge University Press.
- Kalnay, E., et al. (1996). The Ncep/Ncar 40-Year Reanalysis Project. Bulletin of the American Meteorological Society **77**(3): 437-471.
- Kasibhatla, P. and American Geophysical Union (2000). Inverse methods in global biogeochemical cycles. Washington, DC, American Geophysical Union.
- Kistler, R., et al. (2001). The NCEP-NCAR 50-year reanalysis: Monthly means CD-ROM and documentation. Bulletin of the American Meteorological Society **82**(2): 247-267.
- Lamarque, J.-F. and P. G. Hess (2003). Model analysis of the temporal and geographical origin of the CO distribution during the TOPSE campaign. J. Geophys. Res. **108**(D4): 8354.
- Law, B. E., et al. (2002). Environmental controls over carbon dioxide and water vapor exchange of terrestrial vegetation. Agricultural & Forest Meteorology **113**(1-4): 97-120.
- Levin, I., et al. (2002). Observations of atmospheric variability and soil exhalation rate of radon-222 at a Russian forest site. Technical approach and deployment for boundary layer studies. Tellus B **54**(5): 462-475.
- Lloyd, J., et al. (2001). Vertical profiles, boundary layer budgets, and regional flux estimates for CO₂ and its C-13/C-12 ratio and for water vapor above a forest/bog mosaic in central Siberia. Global Biogeochemical Cycles **15**(2): 267-284.
- Olsen, S. C. and J. T. Randerson (2004). Differences between surface and column atmospheric CO₂ and implications for carbon cycle research - art. no. D02301. Journal of Geophysical Research-Atmospheres **109**(D2): 2301.
- Potosnak, M. J., S. C. Wofsy, A. S. Denning, T. J. Conway, J. W. Munger and D. H. Barnes (1999). Influence of biotic exchange and combustion sources on atmospheric CO₂ concentrations in New England from observations at a forest flux tower. Journal of Geophysical Research-Atmospheres **104**(D8): 9561-9569.
- Riley, W. J., J. T. Randerson, P. N. Foster and T. J. Lueker (2004). The Impact of Coastal Meteorology on CO₂ Net Ecosystem Exchange Estimates: Implications for global inversion studies. AGU December, San Francisco, CA.
- Riley, W. J., C. J. Still, B. R. Helliker, M. Ribas-Carbo and J. A. Berry (2003). ¹⁸O composition of CO₂ and H₂O ecosystem pools and fluxes in a tallgrass prairie: Simulations and comparisons to measurements. Global Change Biology **9**(11): 1567-1581.
- Wilkinson, R., K. Clarke, M. Goodchild and J. Reichman (2002). Preparing for Changing Climate: The Potential Consequences of Climate Variability and Change, U.S. Global Change Research Program.
- Winrock (2004). Baseline Report for Forest, Range, and Agricultural Lands of California. Sacramento, California Energy Commission.
- Wofsy, S. C. and R. C. Harris (2002). The North American Carbon Program (NACP), Report of the NACP Committee of the U.S. Interagency Carbon Cycle Science Program., Washington, DC: US Global Change Research Program.
- Xu, L. K. and D. D. Baldocchi (2004). Seasonal variation in carbon dioxide exchange over a Mediterranean annual grassland in California. Agricultural & Forest Meteorology **123**(1-2): 79-96.

Glossary

ABL	Atmospheric Boundary Layer
CARB	California Air Resources Board
CMAQ	Community Multiscale Air Quality Model
CO	Carbon Monoxide
CO ₂	Carbon dioxide
EPA	U.S. Environmental Protection Agency
GHG	Green House Gas
IPCC	Intergovernmental Panel on Climate Change
LADCO	Lake Michigan Air Director Consortium
NACP	North American Carbon Program
NCEP	National Center for Environmental Prediction
NEE	Net Ecosystem Exchange
NEI	EPA National Emission Inventory
NOAA CMDL	U.S. National Oceanic and Atmospheric Administration Climate Monitoring and Diagnostics Laboratory
OCO	NASA Orbiting Carbon Observatory
PIER	Public Interest Energy Research

Appendix A. Diurnal and Seasonal Variations in NEE

Figure A1 shows the diurnal variation in NEE for each ecosystem type in June. Black points indicate the average from all pixels in each day while the red bar indicates the mean (length indicating and standard error of the mean) over the 12 days included in the average. With the exception of (dryland) shrub, the ecosystems take up significant amounts of carbon during the day because temperatures are reasonably warm and water is not seriously limiting and release carbon back the atmosphere at night. The seasonal variation in NEE for the four months included in the simulations are shown in figure A2. Forest, woody savanna, grassland, and shrub ecosystems show seasonal patterns of NEE a combination of sunlight, temperature and moisture responses but are considered to be near carbon balance on an annual basis. Agriculture is modeled as a significant net sink of carbon reflecting the active removal of biomass. Irrigated agriculture is considered uniformly well watered and hence does not experience any water limitation.

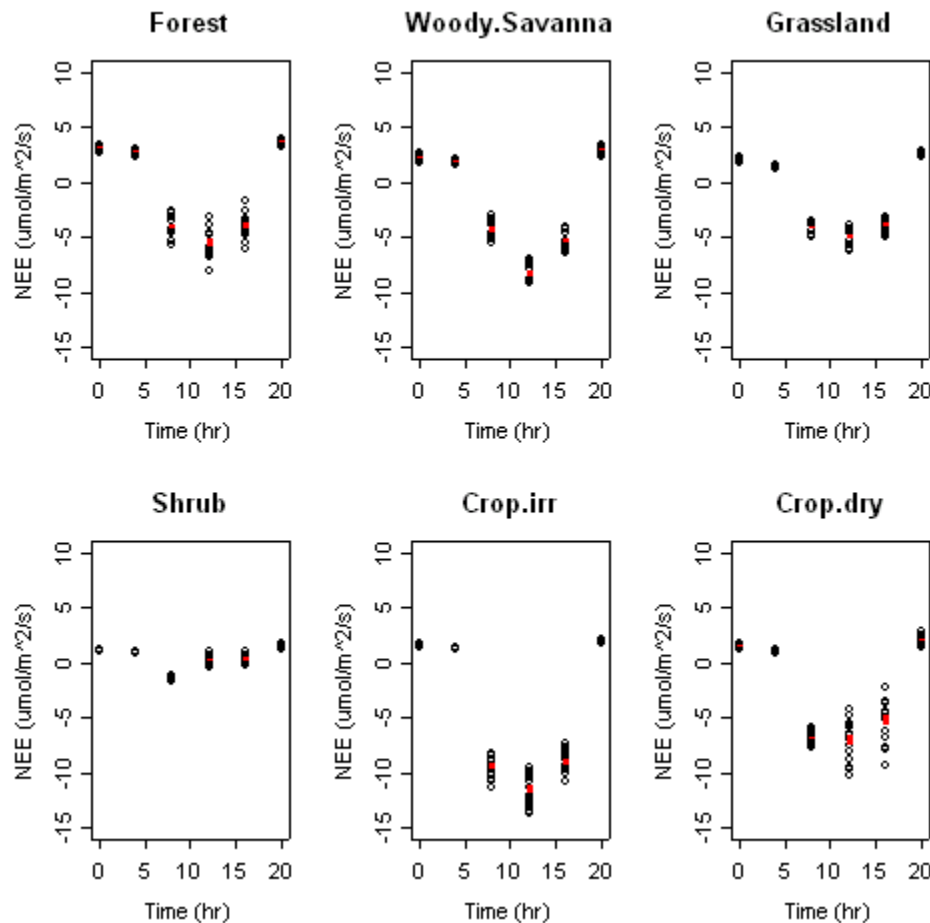


Figure A1 Diurnal variation in June NEE.

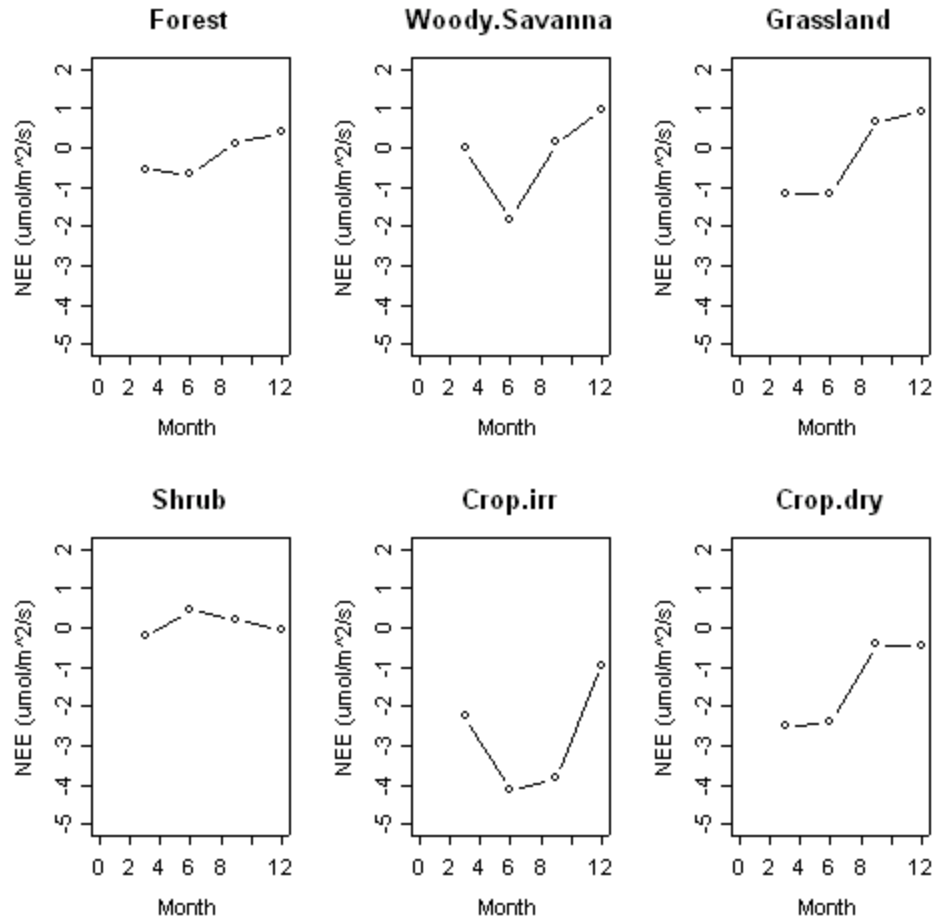


Figure A2. Seasonal variation in daily mean NEE.

Appendix B. Seasonal Cycle of Mean Midday CO₂ Concentration Signals

These figures show the mean midday surface CO₂ concentration signals from NEE inside CA (B1), NEE outside CA (B2), fossil fuel CO₂ inside CA (B3), and fossil fuel CO₂ outside CA (B4). Signals are deviations from the CO₂ background from outside the modeling domain. CO₂ signals from NEE show strong seasonality signals due to ecosystem exchange, while signals from fossil fuel CO₂ show the direction of the mean wind (e.g. predominantly from the north in March), and the amount of mixing relative to average flow. In general, signals from NEE inside CA are large enough to be measured with standard instrumentation at many terrestrial locations.

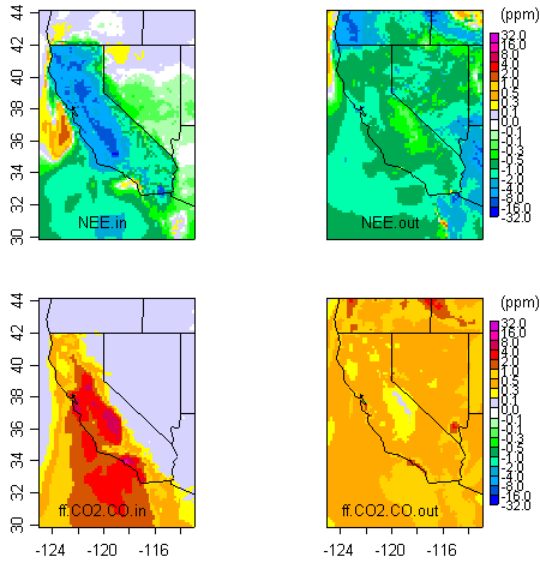


Fig B1. Signals in March.

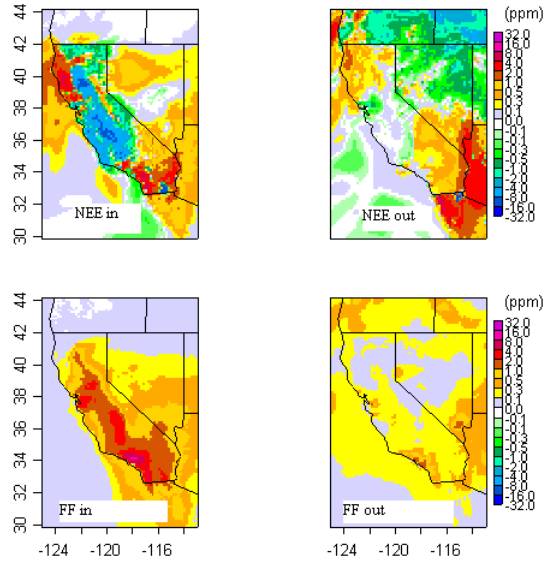


Fig B2. Signals in June (as in Fig6).

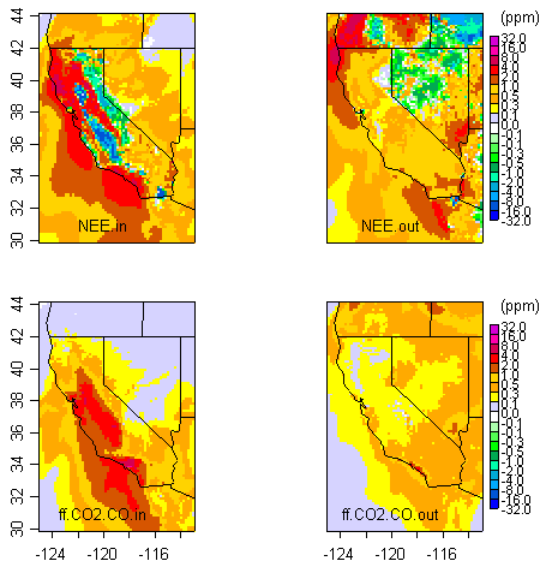


Fig B3. Signals in September.

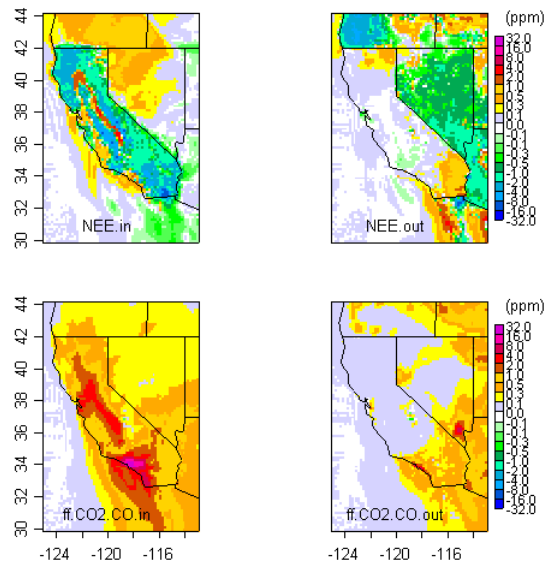


Fig B4. Signals in December.

Appendix C. Seasonal Cycles of Mean Midday Concentration Signals at Observing Stations

The seasonal cycle of mean midday concentration CO₂ signal for each source and observing station are shown in the figure below. For each observing station, the seasonal cycle in CO₂ signal is shown by lines connecting symbols with a common station number. Seasons are ordered March, June, September, December. The standard error of each mean midday value is shown as a thick vertical line. Signals from NEE.in show large seasonal variations because NEE varies with ecology, climate, and atmospheric transport. Signals from fossil fuel combustion show smaller seasonal variations that are caused largely by variations in atmospheric transport.

